

**PALEOENVIRONMENT DATA AND VEGETATION HISTORY FROM A  
SMALL MESOTROPHIC SITE IN THE CURVATURE SUBCARPATHIANS.  
CASE STUDY: INK QUAKING BOG, ROMANIA**

DOI: <http://dx.doi.org/10.18509/GBP.2018.09>

UDC: 551.7(498)

**Vasilică Istrate**<sup>1</sup>

**Alin Mișu-Pintilie**<sup>2</sup>

**Angela Lupașcu**<sup>1</sup>

**Irka Hajdas**<sup>3</sup>

**Emilian Teleaga**<sup>4,5</sup>

<sup>1</sup> Alexandru Ioan Cuza University of Iași, Faculty of Geography and Geology, Department of Geography, Carol I 20A, 700505, Iasi, **Romania**

<sup>2</sup> Alexandru Ioan Cuza University of Iasi, Interdisciplinary Research Department – Field Science, 54 Lascăr Catargiu St., 700107, Iasi, **Romania**

<sup>3</sup> ETH Zurich, Otto-Stern-Weg 58093 Zürich, **Switzerland**

<sup>4</sup> Philipps-Universität Marburg, Biegenstraße 10 35032 Marburg, **Germany**

<sup>5</sup> Vasile Pârvan Archaeological Institute of the Romanian Academy, 11 Henri Coandă str. 010667, Bucharest, **Romania**

**ABSTRACT**

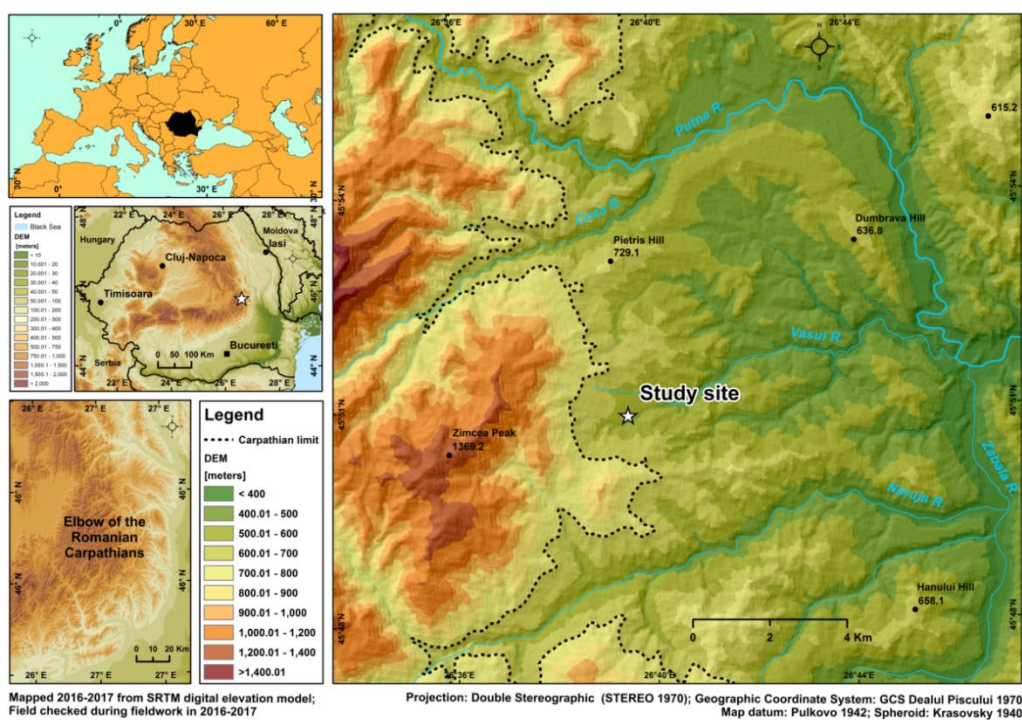
Clearing down the chronology of Holocene's history vegetation was made by exploring some important sites from Romania. The sequences of the forest phases in Holocene are well known because of the studies made by the School of Palynology from Cluj (Romania). These have shown that past vegetation dynamics are not uniform in the Romanian Carpathians, as initially believed. The Ink quaking bog is situated in the contact area of the Carpathians and the Subcarpathians Curvature's sector at the altitude of 560 meters. The surrounding vegetation is highlighted by deciduous forest and meadow or swamp vegetation, having mesotrophic and meso-eutrophic characteristics. The peat bog deposit layer is 70-80 cm thick, and it's represented by a very darkened made of soil peat bog, with some maceration variations. The swamp was dug with a gravity corer, attached with collecting tubes, which allowed a continuous circulation of the inner material column. The samples were collected for the sporopollenin analysis and radiocarbon dating (<sup>14</sup>C). The material collected from the -65 cm layer, dating from 7861±50 B.P., was placed at the limit between Boreal and Atlantic (*Alnus* - 27,6%; *Ulmus* - 16,35%; *Tilia* - 12,78%; *Quercus* - 5,45%; *Picea* - 4,94%; *Pinus* - 4,1%;). The -55 cm layer, dating from 2986±50 B.P., being on the crossover of Subboreal 2 and Subboreal 3 (*Picea* - 11,3%; *Pinus* - 3,91%; *Alnus* - 37,82%; *Ulmus* - 9,56%; *Tilia* - 6,95%). The -35 cm layer, dating from 658±50 B.P., was placed in Subatlantic, were the climate's cooling and rising humidity determined the beech and fir expansion. Based on Ink quaking bog sporopollenin analysis and <sup>14</sup>C dating, we better understand the paleoenvironmental condition around Bîrsești archaeological site (Late Hallstatt).

**Keywords:** vegetation history, radiocarbon dating, mesotrophic quaking bog, Curvature Subcarpathians, Romania

## INTRODUCTION

In Romania, establishing the chronology of Holocene vegetation history was done by investigating key sites in the Carpathian and Sub-Carpathian areas. The succession of the Holocene forest phases is well-known due to the palinological researches developed in the high mountain and sub-mountain peatlands [1], [2], [3], [4], [5], [6], [7], [8], [9]. The current spreading and zoning of flora and vegetation is the result of palaeogeographical and climatic changes that have influenced the succession of the particularities of the cenosis and of the vegetal carpet as a whole [3]. The post-glacial cenogenetic evolution of the Carpathians in Romania differs substantially from the Northern Carpathians and Central Europe. This phenomenon is due to different intensity of glaciation [1], [2], [3], [4], [5]. The territory located on the exterior of the Carpathian arch lies at the contact between the latitudinal vegetation areas characteristic of the Eurasian steppe but also the altitude characteristic of Central Europe [6], [7], [8], [9]. The contact between these areas, more exactly the problem of the relationship between the forest area and the steppe area, has been concerned with many naturalistic researchers, especially since the limit has migrated during the Holocene.

The detailed reconstruction of the history of Holocene vegetation in the Vrancea's Subcarpathians is difficult to achieve because the eutrophic marshes, less acid and with intense microbial activity, do not offer the best conditions for pollen conservation [11], [12], [13]. However, the analysis of peat samples taken from a mesotrophic site may bring new contributions to the knowledge of the environmental conditions of prehistoric civilizations that have habitat in these areas [6]. This study is based on Ink quaking bog (Subcarpathians Curvature's) sporopollenin and radiocarbon dating, and the analysis has been achieved to understand the paleoenvironmental condition and vegetation history around Late Hallstatt Bîrsești archaeological site [10] (Fig. 1).



**Figure 1.** Location map of the studied area in the Curvature Subcarpathians (Romania)

## CASE STUDY: INK QUAKING BOG (CURVATURE SUBCARPATHIANS)

The Ink quaking bog is situated in the contact area of the Carpathians and the Subcarpathians Curvature's sector at the altitude of 560 meters (Fig. 2). The surrounding vegetation is highlighted by deciduous forest and meadow or swamp vegetation, having mesotrophic and meso-eutrophic characteristics. The peat bog deposit layer is 70-80 cm thick, and it's represented by a very darkened made of soil peat bog, with some maceration variations. In the past centuries, artists used natural pigments that they could produce from plants or humic materials. For example, in England and Germany, the Vandyke Brown color refers to the humic pigment taken from the lower layers of the peatland near Cologne [15]. In the villages of the Vrancea's Subcarpathians for the dyeing of wool and fabrics were used certain parts of plants or even turbid material. Ink quaking peat material was used to obtain the black pigment or as a mordant for color fixation. Thus, the name of the analyzed site was attributed by local communities because of the peat's ability to dye some fabrics, of which the most commonly worked was wool.

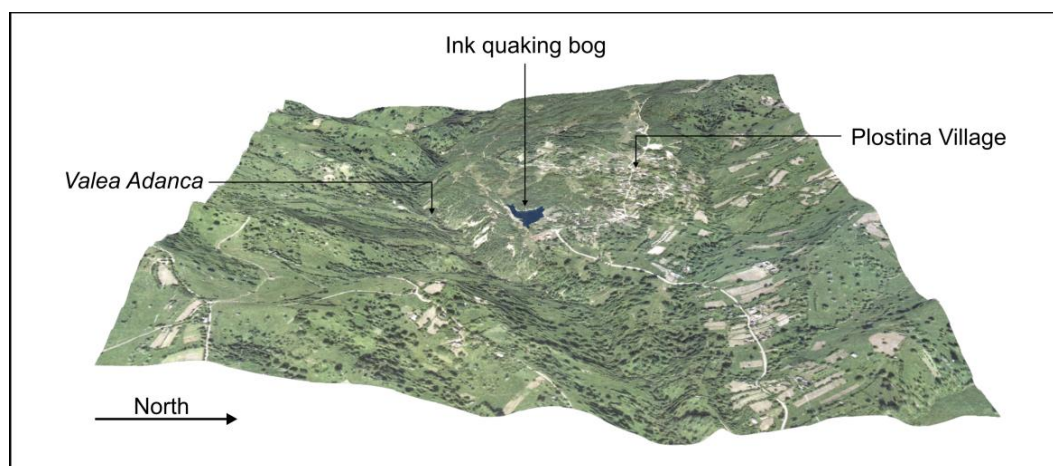


Figure 2. 3D view of the Ink quaking bog and the surrounding area

## METHODS AND TECHNIQUES

The samples were taken for sporo-palynological analysis and for the dating of the radiocarbon dating ( $^{14}\text{C}$ ). The peat layer was drilled using a gravity probe with tubes that allowed the transport of a continuous column of material (10 cm in diameter) into the laboratory. For the sporopollen material analysis samples were taken from 10 in 10 cm, then from 5 in 5 cm. Following the separation of pollen using the Erdtman method, hydrocarbons appeared in the form of perfectly colorless oily droplets. This fact forced us to repeat samples without acetolysis. Palynological preparations were obtained by boiling in 10% KOH, separation by centrifugation and collecting in anhydrous glycerol. We have identified more than 400 tree pollen grains for each sample. Timeline setting was performed sequentially by radiocarbon dating in the ETH Zürich Laboratory.

## RESULTS AND DISCUSSION

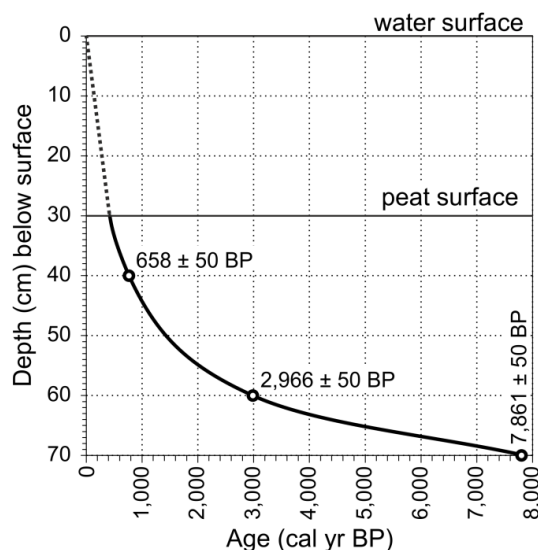
**Chronology.** Of the 4 samples tested for chronology, only 3 were valid: ETH-70053 (depth 30-40 cm) –  $765 \pm 21$  yr BP; ETH-70056 (depth 50-60 cm) –  $2,986 \pm 21$  yr BP; ETH-70057 (depth 60-70 cm) –  $7,861 \pm 24$  yr BP. The sample results from the 40-50 cm depth range (ETH-70055) were not used due to the anomalies most likely caused by the

surface layer mixture at the time of sampling (Table 1). The calibration curve indicates a low sedimentation rate because in only 40 cm accumulated organic material dated in the range 7800-650 BP (Fig. 3).

**Table 1.** ETH  $^{14}\text{C}$  dates from Ink quaking bog

Depth (cm)	Lab. number*	Tested material	$^{14}\text{C}$ yr BP	$F^{14}\text{C}$	Date use (cal yr BP)	Remarks
0 – 30						Water pocket
30 – 40	ETH-70053	Macrofossils	765±21	0.9092	765	no remarks
40 – 50	ETH-70055	Macrofossils	658±21	0.9213		anomalous not used
50 – 60	ETH-70056	Wood	2,986±21	0.6896	2,986	no remarks
60 – 70	ETH-70057	Macrofossils	7,861±24	0.3758	7,861	no remarks

\*ETH Zürich Laboratory;  $^{14}\text{C}$  age [BP (before 1950 AD)]; Calendar age  $2\sigma$  range AD/BC -- calibrated ranges 95% con. Level: Fraction Modern:  $F^{14}\text{C} = \exp(-F^{14}\text{C age}/8033)$ ; If  $F^{14}\text{C} > 1$ , the sample indicates presence of „bomb peak  $^{14}\text{C}$ ” (post 1950AD)



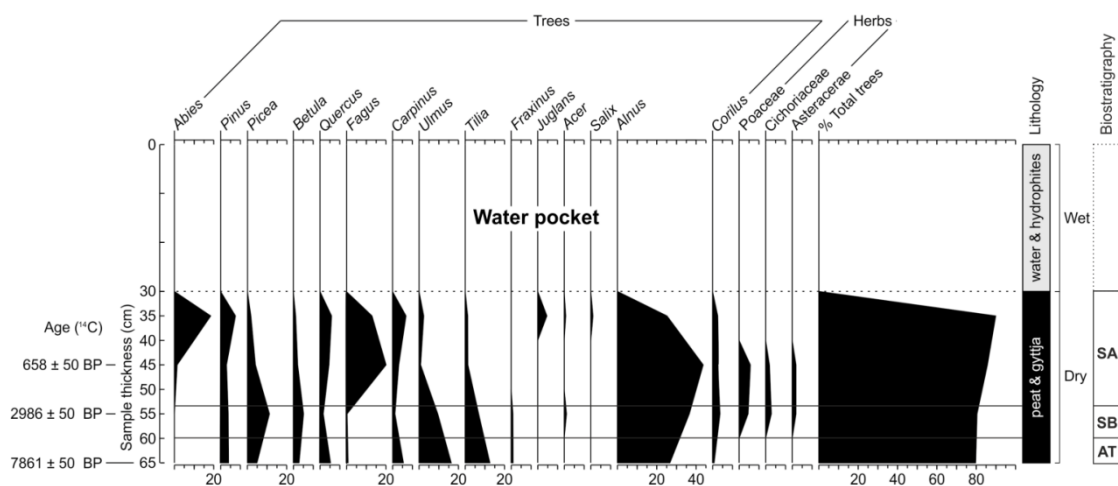
**Figure 3.** Calibrated radiocarbon data and age–depth model for Ink quaking bog

**Vegetation history.** The material collected from a depth of 65 cm ( $7,861 \pm 24$  yr BP) was deposited at the Boreal-Atlantic limit. The forests were predominantly of *Ulmus* (16.35%), *Tilia* (12.78%) and *Quercus* (5.45%) in much *Alnus* (27.6%). Other deciduous genres were present in 1-2%. Among the conifers, *Picea* (4.94%) show an expansion trend and *Pinus* (4.09%) reduction, being approximately equally represented. The herbaceous layer was richer and more diversified than in the other analyzed samples. The variation of the curves recorded on the thermophilic elements is less pronounced as a result of the lower pollen productivity of these essences on the one hand and on the other hand due to the lower expansion of the thermophilic elements in the Oriental Carpathians to the Apuseni Mountains [3], [4], [5]. The fluctuations of herbage pollen curves also highlight a particularity of the climate during this period as a result of increased humidity due to higher temperatures that have been demonstrated since Boreal and continued in the Atlantic and Subboreal [1], [2], [3] (Fig. 4).

The material collected from the depth of 55 cm deep ( $2,986 \pm 21$  yr BP) was deposited at the limit between Subboreal 2 (Sb2) and Subboreal 3 (Sb3) and is not very rich in pollen. In some studies, for polinic spectrum in the Prut river,  $3290 \pm 80$  yr BP for Sb2, and 2650

$\pm 50$  yr BP for Sb3 [9]. Another argument that supports the placement of our probe at the Sb2/Sb3's limit is that of a low frequency given by the *Carpinus*, with 3%, knowing that this species had 2 pic-uri in Sb1 (*Carpinus betulus* – species found on the plains or hills, being resistant to cold and late freezing, but vulnerable to early heat, drought or dryness) and Sb3 (*Carpinus orientalis* – thermophilic species, very xerophytic, not pretentious to the soil, flourishing in a warm environment) and a redundancy between the two (our case scenario).

Higher values are found for *Picea* (11.3%) and *Pinus* (3.91%), and *Abies* only present. Among the most represented species were *Alnus* (37.82%), *Ulmus* and *Tilia* (9.56% and 6.95% respectively) and even *Betula* and *Corylus* 4-5%. Elm and lime are thermophilic species, pretentious to climate and soil; the elm is eutrophic and eurifit species with great valences to adapt to soil moisture. Among the grass species belonging to the Poaceae families, Cichoriaceae and Asteraceae were more common (2-5%). The climate during this period was warm and with a high degree of dryness, revealed by the xerophilic herbaceous species. For the same period, the analyzes in the Bisoca marsh [3] shows the same decrease of the hornbeam compared to a previous period and the continued presence of the pottage (*Plantago lanceolata*). This dynamic suggests anthropogenic pressure on the environment, also evidenced by the presence of rye pollen in the period 2500-1200 yr BP [3] (Fig. 4).



**Figure 4.** Pollen diagrams from Ink quaking bog: AT – Atlantic, SB – Subboreal, SA – Subatlantic

The 45-cm-deep sample, dating from  $658 \pm 50$  B.P. years, entered the Subatlantic time. The climate's cooling and rising humidity were factors that determined the beech and fir expansion. The poline obtained is rich in pollen grains of trees, shrubs and grasses. The polar spectrum reveals predominance of foies, the alder shielding almost the other hardwoods: *Alnus* 5.08%. Conifers were represented by the three genres: *Picea* 4.1%, *Pinus* 3.07% and *Abies* 1.43%. The best-represented species, apart from the alder that is present throughout Holocene, is beagle (*Fagus* 20.08%) and which certifies the period we are in. Beech vegetates in a continental mountain climate, being demanding for humidity and precipitation. It is also sensitive to drought, dryness, late frosts and early frosts as well as excessive frosts. Favorable environmental conditions for beech and fir have become less favorable for lime (1.43%) and elm (0.82%) which have been greatly reduced. It is also sensitive to drought, dryness, late frosts and early frosts as well as

excessive frosts. Favorable environmental conditions for beech and fir have become less favorable for lime (1.43%) and elm (0.82%) which have been greatly reduced. (Fig. 4).

**Connection with paleoenvironment and prehistoric civilizations.** Chronologically, the Iron Age occurred during the last two periods of the Post-glacier: Sub-Boreal and Subatlantic. These two periods are characterized by climate change that has greatly influenced the way of life of human populations [16].

Chronologically, the Iron Age existed between the last couple post-glacial' periods, known as the Subboreal and Subatlantic. These two periods are characterized by climate change which greatly influenced the human's way of living [16]. Based on sporopollenic analyzes and the results of other studies with similar themes [6], we conclude that on Hallstatt's passage the climatic variability directly influenced the way of life of prehistoric populations. The first part of the Subboreal (4000-3200 yr BP) characterized by a warmer and more dry climate is the period when the first cereal spores occurred in the Carpathian Bend area as a sign of agricultural activity [3].

After Ink Lake's the sporopollenic analysis and also because of other studies [6], we can conclude by saying that over the Hallstatt, climate's variability directly influenced the way of how prehistoric men used to live. The first half of the Subboreal (4000-3200 B.P.) characterized by a warmer and dryer climate, was the first time when methods of yield enhancement for cereal came into sight in the Curvature Carpathians area as a sign of agricultural activity [3]. The anthropogenic relief forms represented by the agroterages [14] on the right side of Putna river (Bîrsești area), whose age is not yet determined, are definitely worth mentioning, but they are certainly preceded by the first historical sources that refer to this area.

The next two parts of the Subboreal are characterized by climatic conditions similar to those of the current period with a cooling at the passage between Sb2 / Sb3. In this range between 3200-2600 yr BP, the presence of plants such as *Plantago lanceolata* suggests an anthropic pressure exerted mainly by grazing. The transition to the Sub-Atlantic (2600 yr BP-present) has brought a cooler and more humid climate with the extension of the beech forests, but agricultural practices are still reported through the presence of rye pollen.

## CONCLUSIONS

The Ink quaking bog is a small mesotrophic site situated in the contact area of the Carpathians and the Subcarpathians Curvature's (Romania). Radiocarbon dating indicated a maximum age of marshes of  $7.861 \pm 24$  yr BP. Sporopollenic analyzes indicate that climatic variability directly influenced the way of life of prehistoric populations on Hallstatt. In Sb1 (4000-3200 yr BP) is the period when the first cereal spores appear in the Carpathian Bend area as a sign of agricultural activity. In Sb2 / Sb3 (3200-2600 yr BP), the presence of plants such as *Plantago lanceolata* suggests anthropogenic pressure exerted by pasturage. In the Subatlantic (2600 yr BP - present) there are still signs of practicing agriculture through the presence of rye pollen. The results obtained in this study contribute to the reconstruction of the paleoenvironment of the archaeological site on the Dumbrava Hill (Bîrsești), 8 km from Ink quaking bog.

## ACKNOWLEDGMENTS

This work was financially supported by the project PN-II-RU-TE-2014-4-1602, "Late Hallstatt burials and funerary practices at the Lower Danube" and the Department of

Geography from the “Alexandru Ioan Cuza” University of Iasi, Romania, and the infrastructure was provided through the POSCCE-O 2.2.1, SMIS-CSNR 13984-901, No. 257/28.09.2010 Project, CERNESIM (L4) and Geoarchaeology Laboratory.

## REFERENCES

- [1] Tantau I., Reille M., deBeaulieu J.L., Farcas S., Goslar T. & Paterne M. Vegetation history in the eastern Romanian Carpathians: pollen analysis of two sequences from the Mohos crater, *Vegetation History and Archaeobotany*, vol. 12(2), pp 113-125, 2003.
- [2] Tantau I., Reille M., deBeaulieu J.L. & Farcas S. Late Glacial and Holocene vegetation history in the southern part of Transylvania (Romania): pollen analysis of two sequences from Avrig, *Journal of Quaternary Science*, vol. 21(1), pp 49-61, 2006.
- [3] Tantau I., Reille M., deBeaulieu J.L. & Farcas S. Holocene vegetation history in Romanian Subcarpathians, *Quaternary Research*, vol. 72(2), pp 164-173, 2009.
- [4] Rösch M. & Fischer E. A radiocarbon dated Holocene pollen profile from the Banat mountains (Southwestern Carpathians, Romania), *Flora*, vol. 195, pp 277-286, 2000.
- [5] Feurdean A., Wohlfarth B., Björkman S., Tantau I., Bennike O., Willis K.J., Farcas S. & Robertsson A.M. The influence of refugial population on Lateglacial and early Holocene vegetational changes in Romania, *Review of Palaeobotany and Palynology*, vol. 145, pp 305-320, 2007.
- [6] Feurdean A. Holocene forest dynamics in northwestern Romania, *The Holocene*, vol. 13, pp 435-446, 2005.
- [7] Feurdean A. & Bennike O. Late Quaternary palaeoecological and paleoclimatological reconstruction in the Gutaiului Mountains, NW Romania, *Journal of Quaternary Science*, vol. 19, pp 809-827, 2004.
- [8] Björkman S., Feurdean A., Cinthio K., Wohlfarth B. & Possnert G. Late Glacial and early Holocene vegetation development in the Gutâiului Mountains, northwestern Romania, *Quaternary Science Reviews*, vol. 21, pp 1039–1059, 2002.
- [9] Gałka M., Tantau I., Ersek V. & Feurdean A. A 9000 year record of cyclic vegetation changes identified in a montane peatland deposit located in the Eastern Carpathians (Central-Eastern Europe): Autogenic succession or regional climatic influences?, *Palaeogeography, Palaeoclimatology, Palaeoecology*, vol. 449, pp 52-61, 2016.
- [10] Teleagă E., Sârbu D., Constantinescu M., Stoica S., Istrate V. Neue archäologische untersuchungen derhallstattzeitlichen nekropole in Bârsești, jud. Vrancea, *Caiete ARA*, vol.6, 2015, pp. 151-166, 2015.
- [11] Mișu-Pintilie A., Asandulesei A., Nicu I.C., Stoleriu C.C. & Romanescu G. Using GPR for assessing the volume of sediments from the largest natural dam lake of the Eastern Carpathians: Cujdel Lake, Romania, *Environmental Earth Sciences*, vol. 75(8), No. 710, 2016. Doi: 10.1007/s12665-016-5537-1
- [12] Mișu-Pintilie A., Paiu M., Breaban I.G. & Romanescu G. Status of water quality in Cujdi hydrographic basin from Eastern Carpathian, Romania, *SGEM2014 Conference Proceedings, Geoconference on Water Resources, Forest, Marine and Ocean Ecosystems, Bulgaria*, vol. 1, pp 639-646, 2014.
- [13] Cozma D.G., Cruceanu A., Cojoc G.M., Muntele I. & Mișu-Pintilie A. The factorial analysis of physico-chemical indicators in Bistrita's upper hydrographical basin, *SGEM2015 Conference Proceedings, Geoconference on Water Resources, Forest, Marine and Ocean Ecosystems, Bulgaria*, vol. 1, pp 625-632, 2015.
- [14] Ursu A., Chelaru D. A., Mihai F. C., Iordache I., 2011, Anthropogenic landform modeling using GIS techniques case study: Vrancea region, *Geographia Technica*, vol. 6, Issue 1, pp. 91-100, 2011.
- [15] Jonhston-Feller R. *Color Science in the Examination of Museum Objects: Nondestructive Procedures*, The Getty Conservation Institute, SUA-Los Angeles, 2001.

- [16] Renssena H., Goosse H. & Fichefet T. Simulation of Holocene cooling events in a coupled climate model, *Quaternary Science Reviews*, vol. 26, pp 2019–2029, 2007.